Geophysical Research Abstracts, Vol. 10, EGU2008-A-09749, 2008 SRef-ID: 1607-7962/gra/EGU2008-A-09749 EGU General Assembly 2008 © Author(s) 2008



Volatiles in kimberlite: Volume relationships and implications for conduit and eruption dynamics

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The eruption of kimberlite magma must be governed by the same physical and chemical principles controlling eruption of basaltic to rhyolitic magmas. However, this does not mean that kimberlite behavior can be understood by simple linear extrapolation of ideas developed for other volcanoes. The most unique aspect of kimberlite magmas is their potential for having high dissolved contents of primary volatiles (e.g., H_2O + > 15 wt. %) coupled to a high ascent rate. The high ascent rates ensures that CO_2 higher volatile contents are maintained to the point of eruption. Here, we use recent results in experimental and theoretical petrology [1, 2] to explore and amplify concepts that have been advanced for the emplacement and eruption of kimberlite [3, 4, 5, 6]. Specifically, recently published thermodynamic models for H₂O-CO₂-silicate melt equilibria [1, 2] are used to simulate melt-solid-fluid equilibria during emplacement and eruption of kimberlite. These new thermodynamic models forecast volatile solubilities as a function of T, P and melt composition and can track energetics (e.g., enthalpy, heat capacity) and volume relationships. Heat contents are critical for evaluating the extent to which these systems are able to cool during eruption relative to their characteristic glass transition temperatures [7, 8, 9]. Efficient cooling can promote glass formation. More importantly, these calculations can explore the volume relationships between melt, fluid and solid [4, 7, 10, 11]. These relationships have direct consequences for the style and duration of volcanic eruption, the size and stability of the volcanic plume, depths of magmatic fragmentation, and conduit evolution. We use these calculations to test the premise that the geometries of the Diavik kimberlite pipes (Diameter: 120-140 m; Depth: 350-400 m) are ultimately a reflection of eruption intensity (e.g., diameter α volume flux) and duration (depth α time).

References Cited:

[1] Nicholls, J 2000. Thermodynamics of a Magmatic gas phase" 50 years later: Comments on a paper by JohnVerhoogen (1949). Can Mineral, 38: 1313-1328.

[2] Papale P 1999. Modeling of the solubility of a two-component H_2O+CO_2 fluid in silicate liquids. Am Mineral 84: 477-492.

[3] Scott Smith, BH & Field, M. 1999. Contrasting geology and near-surface emplacement of kimberlite pipes in southern Africa and Canada, 7th Int Kimberlite Conf Abstr

[4] Kjarskaard, B. 2003. Behaviour of kimberlite magma in the upper crust and at surface. 8th Int Kimberlite Conf Abstr

[5] Wilson, L & Head, JW 2003. Diatremes and Kimberlites 2: An integrated model of the ascent and eruption of kimberlitic magmas and the production of crater, diatreme, and hypabyssal facies. 8th Int Kimberlite Conf Abstr

[6] Spera, F.J. 1984. Carbon dioxide in petrogenesis III: role of volatiles in the ascent of alkaline magma with special reference to xenolith-bearing mafic lavas. Contrib Mineral Petrol, 88: 217-232.

[7] Sparks, R.S.J. 1978. The dynamics of bubble formation and growth in magmas: a review and analysis. J Volcanol Geotherm Res 3, 1-37.

[8] Nicholls, J & Stout, MZ 1982. Heat effects of assimilation, crystallization, and vesiculation in magmas. Contrib Mineral Petrol 81: 328-339.

[9] Russell, JK 1987. Crystallization and vesiculation of the 1984 Eruption of Mauna Loa. J Geophys Res 92:13731-13743.