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## Oxygen three-isotope distributions in tropospheric and stratospheric $CO_2$ – potential inferences from high precision isotope measurements of carbonate decomposition

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It was demonstrated several years ago (Miller et al., *Proc. Natnl. Acad. Sci. USA* **99**, 10988–10993, 2002) that thermal decomposition of carbonates (including NBS18 and NBS19), under high vacuum conditions to minimize back-reaction, is accompanied – remarkably – by a small but distinctive 'mass-independent' isotopic fractionation, with the resulting CO<sub>2</sub> being anomalously enriched in <sup>17</sup>O ( $\Delta^{17}O \sim 0.15$  per mil). Re-examination of isotopic data from both CO<sub>2</sub> and solid oxide reaction products indicates that terrestrial carbonates appear to fit a fractionation line of slope 0.524 on the logarithmic form of the oxygen three-isotope plot. This is notably lower than the value of 0.528 which characterises natural waters. Furthermore, the few published high precision measurements of both  $\delta^{17}$ O and  $\delta^{18}$ O in tropospheric CO<sub>2</sub> suggest that its oxygen three-isotope composition is also in accord with that of carbonate minerals, rather than waters. This finding is of relevance to the '<sup>17</sup>O correction' to  $\delta^{13}$ C measurements of CO<sub>2</sub> at the highest levels of accuracy and precision.

In contrast, CO<sub>2</sub> from the lower and middle stratosphere is characterised by a slope of 1.7 on the oxygen three-isotope plot (Lämmerzahl et al., *Geophys. Res. Lett.*, **29**, 1582, 10.1029/2001GL014343, 2002). Whereas CO<sub>2</sub> + O(<sup>1</sup>D)  $\rightarrow$  CO<sub>3</sub><sup>\*</sup>  $\rightarrow$  CO<sub>2</sub> +

 $O(^{3}P)$  is generally regarded as the mechanism by which the  $^{17}O$  and  $^{18}O$  enrichments in stratospheric ozone are transferred into coexisting CO<sub>2</sub>, a satisfactory explanation for the slope of the fractionation line remains elusive. The possibility that formation and/or decomposition of the CO3\* might also be associated with a 'mass independent' isotopic fractionation was first suggested by Wen and Thiemens (J. Geophys. Res. 98 D7, 12801–12808, 1993). If thermal decomposition of the carbonate ion is analogous to the  $CO_3^* \rightarrow CO_2 + O(^{3}P)$  reaction, isotopic behaviour during the former may be relevant to an explanation of the very unusual oxygen three-isotope distribution in stratospheric  $CO_2$ . However, whereas the carbonate ion has  $D_{3h}$  symmetry, it has long been known that the ground state  $CO_3^*$  isomer has  $C_{2v}$  symmetry, although interconversion to the  $D_{3h}$  isomer should readily occur, as the energy state of the latter is only 0.42 kJ mol<sup>-1</sup> higher, with an isomerization barrier of 18.4 kJ mol<sup>-1</sup>. The first experimental detection of the  $D_{3h}$  isomer of  $CO_3^*$  occurred only recently (Jamieson et al., COSPAR 2006-A-03495), but this discovery strengthens the potential link with the carbonate decomposition reaction. As the lifetime of  $CO_3^*$  is only  $\sim 10^{-12}$ s, multiple cycles of the formation and decomposition of this entity from  $CO_2 + O(^1D)$  in the stratosphere might be a contributory mechanism to not only the range of <sup>18</sup>O enrichments of stratospheric CO<sub>2</sub> relative to the tropospheric CO<sub>2</sub> reservoir, but also to the attendant slope of 1.7 on the oxygen three-isotope plot.